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6 **HAAR WAVELET ANALYSIS OF CLIMATIC TIME SERIES**

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30 In order to extract the intrinsic information of climatic time series from background red  
31 noise, in this paper, we will first give an analytic formula on the distribution of Haar  
32 wavelet power spectra of red noise in a rigorous statistical framework. After that, by  
33 comparing the difference of wavelet power spectra of real climatic time series and red  
34 noise, we can extract intrinsic features of climatic time series. Finally, we use our method  
35 to analyze Arctic Oscillation (AO) which is a key aspect of climate variability in the  
36 Northern Hemisphere, and discover a great change in fundamental properties of the AO,  
37 commonly called a regime shift or tripping point.

38 *Keywords:* Climatic time series; Haar wavelet analysis; red noise.

39 AMS Subject Classification: 42C

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## 1 Introduction

The continuous wavelet transform possesses the ability to construct a time-frequency representation of a time series that offers very good time and frequency localization, so wavelet transforms can analyze localized intermittent periodicities of climatic time series.<sup>1,7,8</sup> The relative resolution in time and frequency space depends on the wavelet chosen. Many climatic time series have been analyzed with the Morlet wavelet.<sup>2,3,5</sup> which is roughly equally localized in time and frequency. However, the Haar wavelet which we discuss has not been previously used in this respect.

Physical interpretation of climatic time series requires statistical significant testing against a null hypothesis of some noise model. Different from regular signal analysis, climatic time series analysis often use red noise as noise model.<sup>4</sup> A simple model for red noise is the univariate lag-1 autoregressive (AR(1)) process. By comparing the difference of wavelet power spectra of real climatic series and AR(1) red noise, one can extract the intrinsic features of climatic time series. To date significance testing has been based on an empirical formula for the Morlet wavelet power spectra of AR(1) red noise.<sup>2,3,5</sup>

In this paper, we will focus on Haar wavelet. We will give an analytic formula of the distribution of Haar wavelet power spectra for an AR(1) red noise in a rigorous statistical framework (see Sec. 2). The relation between scale  $a$  and Fourier period  $T$  for the Morlet wavelet is  $a = 0.97T$ .<sup>5</sup> However, for Haar wavelet, the corresponding formula is  $a = 0.3710T$  (see Appendix A). Since for any time series of time step  $\delta t$  and total length  $N\delta t$ , the range of scales is from the smallest resolvable scale  $2\delta t$  to the largest scale  $N\delta t$  in wavelet-based time series analysis, by using the Haar wavelet analysis, one can extract more low frequency intrinsic information. At the end of this paper, we use Haar wavelet to do the statistical significance test of the Arctic Oscillation (AO) and discover a great change in fundamental properties of the AO, commonly called a regime shift or tripping point.

## 2 Haar Wavelet Power Spectrum Distribution for AR(1) Time Series

Different from regular signal analysis, climatic time series analysis often use red noise as noise model.<sup>4</sup> A simple model for red noise is the univariate lag-1 autoregressive (AR(1)) process. A discrete random process  $\{x_n\}_0^{N-1}$  is called an AR(1) time series with the parameter  $\alpha$  if

$$x_n = \alpha x_{n-1} + Z_n, \quad n = 1, \dots, N-1 \quad \text{and} \quad x_0 = 0,$$

where the parameter  $0 \leq \alpha < 1$  and  $\{Z_n\}_1^\infty$  is the white noise, i.e. they are independent normal distributed variables with variance  $\sigma^2$ . From this, we can successively obtain that

$$x_n = \sum_{k=1}^n \alpha^{n-k} Z_k \tag{2.1}$$

and the expectation

$$E[x_n] = \sum_{k=1}^n \alpha^{n-k} E[Z_k] = 0.$$

Therefore, each  $x_n$  is a Gaussian random variable with mean 0. Let  $H(x)$  be the Haar function:

$$H(x) = \begin{cases} 1, & 0 < x \leq \frac{1}{2}, \\ -1, & -\frac{1}{2} < x < 0, \\ 0, & \text{otherwise.} \end{cases} \quad (2.2)$$

The wavelet transform  $W_\nu(s)$  of AR(1) time series  $\{x_n\}_0^{N-1}$  associated with  $H(x)$  is defined ordinarily as<sup>2,3</sup>:

$$\sqrt{\frac{s}{\delta t}} W_\nu(s) = \sum_{n=0}^{N-1} x_n H\left(\frac{(n-\nu)\delta t}{s}\right), \quad (2.3)$$

where  $\nu$  is the time parameter and  $s$  is the scaling parameter and  $\delta t$  is a sample period of  $\{x_n\}_0^{N-1}$ . By (2.2),

$$H\left(\frac{(n-\nu)\delta t}{s}\right) = \begin{cases} 1 & \left(\nu < n \leq \nu + \frac{s}{2\delta t}\right), \\ -1 & \left(\nu - \frac{s}{2\delta t} < n < \nu\right), \\ 0 & \left(n = \nu \text{ or } n \leq \nu - \frac{s}{2\delta t} \text{ or } n > \nu + \frac{s}{2\delta t}\right). \end{cases} \quad (2.4)$$

We choose  $\nu, s$  such that

$$\frac{s}{2\delta t} \leq \nu \leq N - 1 - \frac{s}{2\delta t}. \quad (2.5)$$

By (2.3),

$$\sqrt{\frac{s}{\delta t}} E[W_\nu(s)] = \sum_{n=0}^{N-1} E[x_n] H\left(\frac{(n-\nu)\delta t}{s}\right) = 0.$$

Therefore,  $W_\nu(s)$  is a Gaussian random variable with mean 0. To obtain the distribution of the power spectrum  $|W_\nu(s)|^2$ , we only need to compute the variance of  $W_\nu(s)$ . Note that

$$\text{Var}(W_\nu(s)) = E[(W_\nu(s))^2] - (E[W_\nu(s)])^2 = E[(W_\nu(s))^2].$$

From this and (2.3), we have

$$\begin{aligned} \frac{s}{\delta t} \text{Var}(W_\nu(s)) &= E \left[ \left( \sum_{n=0}^{N-1} x_n H\left(\frac{(n-\nu)\delta t}{s}\right) \right)^2 \right] \\ &= \sum_{n,m=0}^{N-1} E[x_n x_m] H\left(\frac{(n-\nu)\delta t}{s}\right) H\left(\frac{(m-\nu)\delta t}{s}\right). \end{aligned} \quad (2.6)$$

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Denote  $s^* = \frac{s}{2\delta t}$ . By (2.5), we have

$$\nu - s^* \geq 0 \quad \text{and} \quad \nu + s^* \leq N - 1.$$

From this and (2.4), by (2.6), we get

$$2s^* \text{Var}(W_\nu(s)) = \sum_{n,m=\nu-s^*}^{\nu+s^*} E[x_n x_m] H\left(\frac{m-\nu}{2s^*}\right) H\left(\frac{n-\nu}{2s^*}\right). \quad (2.7)$$

Since  $\{Z_n\}_1^\infty$  is white noise with variance  $\sigma^2$ ,

$$E[Z_k Z_l] = 0 (k \neq l) \quad \text{and} \quad E[Z_k^2] = \sigma^2.$$

By (2.1), we get

$$\begin{aligned} E[x_n x_m] &= E \left[ \left( \sum_{k=1}^m \alpha^{m-k} Z_k \right) \left( \sum_{k=1}^n \alpha^{n-k} Z_k \right) \right] \\ &= \sum_{k=1}^{\lambda_{mn}} \alpha^{n-k} \alpha^{m-k} E[Z_k^2] \\ &= \sigma^2 \alpha^{n+m} \sum_{k=1}^{\lambda_{mn}} \alpha^{-2k} = \frac{\sigma^2 \alpha^{n+m} (1 - \alpha^{-2\lambda_{mn}})}{\alpha^2 - 1}, \end{aligned}$$

where  $\lambda_{mn} = \min\{n, m\}$ . From this, we get that

$$\left( \frac{\alpha^2 - 1}{\sigma^2} \right) E[x_n x_m] = \begin{cases} \alpha^{m+n} - \alpha^{n-m}, & m \leq n, \\ \alpha^{m+n} - \alpha^{m-n}, & m > n. \end{cases} \quad (2.8)$$

So, we get

$$\left( \frac{\alpha^2 - 1}{\sigma^2} \right) E[x_n x_m] = \alpha^{m+n} - \alpha^{|m-n|}.$$

From this and (2.7), we have

$$\begin{aligned} -2s^* \left( \frac{\alpha^2 - 1}{\sigma^2} \right) \text{Var}(W_\nu(s)) &= \sum_{n,m=\nu-s^*}^{\nu+s^*} (\alpha^{|m-n|} - \alpha^{m+n}) H\left(\frac{m-\nu}{2s^*}\right) \\ &\quad \times H\left(\frac{n-\nu}{2s^*}\right) = J - \tilde{J}, \end{aligned} \quad (2.9)$$

where

$$J = \sum_{n,m=\nu-s^*}^{\nu+s^*} \alpha^{|m-n|} H\left(\frac{m-\nu}{2s^*}\right) H\left(\frac{n-\nu}{2s^*}\right)$$

and

$$\tilde{J} = \sum_{n,m=\nu-s^*}^{\nu+s^*} \alpha^{m+n} H\left(\frac{m-\nu}{2s^*}\right) H\left(\frac{n-\nu}{2s^*}\right). \quad (2.10)$$

1 We first compute  $J$ .

By (2.4), we divide  $J$  into four sums:

$$J = \left( \sum_{n,m=\nu-s^*}^{\nu-1} - \sum_{m=\nu-s^*}^{\nu-1} \sum_{n=\nu+1}^{\nu+s^*} - \sum_{m=\nu+1}^{\nu+s^*} \sum_{n=\nu-s^*}^{\nu-1} + \sum_{n,m=\nu+1}^{\nu+s^*} \right) \alpha^{|m-n|} \\ =: J_1 - J_2 - J_3 + J_4. \quad (2.11)$$

We compute each sum. For  $J_1$ :

$$J_1 = \left( \sum_{m=\nu-s^*}^{\nu-1} \sum_{n=\nu-s^*}^m + \sum_{m=\nu-s^*}^{\nu-1} \sum_{n=m+1}^{\nu-1} \right) \alpha^{|m-n|} = J_{11} + J_{12},$$

where

$$J_{11} = \sum_{m=\nu-s^*}^{\nu-1} \alpha^m \left( \sum_{n=\nu-s^*}^m \alpha^{-n} \right) \\ = \sum_{m=\nu-s^*}^{\nu-1} \frac{\alpha^m (\alpha^{-\nu+s^*} - \alpha^{-m-1})}{1 - \alpha^{-1}} = \frac{\alpha^{-\nu+s^*}}{1 - \alpha^{-1}} \sum_{m=\nu-s^*}^{\nu-1} \alpha^m - \sum_{m=\nu-s^*}^{\nu-1} \frac{\alpha^{-1}}{1 - \alpha^{-1}} \\ = -\frac{(\alpha^{-\nu+s^*})(\alpha^\nu - \alpha^{\nu-s^*})}{(1 - \alpha)(1 - \alpha^{-1})} - \frac{s^*}{\alpha(1 - \alpha^{-1})} = \frac{1 - \alpha^{s^*}}{(1 - \alpha^{-1})(1 - \alpha)} + \frac{s^*}{1 - \alpha}$$

and

$$J_{12} = \sum_{m=\nu-s^*}^{\nu-1} \alpha^{-m} \left( \sum_{n=m+1}^{\nu-1} \alpha^n \right) \\ = \sum_{m=\nu-s^*}^{\nu-1} \frac{\alpha^{-m} (\alpha^{m+1} - \alpha^\nu)}{1 - \alpha} = -\frac{\alpha^\nu}{1 - \alpha} \sum_{m=\nu-s^*}^{\nu-1} \alpha^{-m} + \frac{1}{1 - \alpha} \sum_{m=\nu-s^*}^{\nu-1} \alpha \\ = -\frac{\alpha^\nu (\alpha^{-\nu+s^*} - \alpha^{-\nu})}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{\alpha s^*}{1 - \alpha} = \frac{1 - \alpha^{s^*}}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{\alpha s^*}{1 - \alpha}.$$

Therefore, we have

$$J_1 = J_{11} + J_{12} = \frac{2(1 - \alpha^{s^*})}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{(1 + \alpha)s^*}{1 - \alpha}.$$

For  $J_4$ :

$$J_4 = \sum_{m=\nu+1}^{\nu+s^*} \left( \sum_{n=\nu+1}^m + \sum_{n=m+1}^{\nu+s^*} \right) \alpha^{|m-n|} = J_{41} + J_{42},$$

where

$$J_{41} = \sum_{m=\nu+1}^{\nu+s^*} \alpha^m \left( \sum_{n=\nu+1}^m \alpha^{-n} \right)$$

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$$\begin{aligned} &= \sum_{m=\nu+1}^{\nu+s^*} \frac{\alpha^m(\alpha^{-\nu-1} - \alpha^{-m-1})}{1 - \alpha^{-1}} = \frac{\alpha^{-\nu-1}}{1 - \alpha^{-1}} \sum_{m=\nu+1}^{\nu+s^*} \alpha^m - \sum_{m=\nu+1}^{\nu+s^*} \frac{\alpha^{-1}}{1 - \alpha^{-1}} \\ &= \frac{\alpha^{-\nu-1}(\alpha^{\nu+1} - \alpha^{\nu+s^*+1})}{(1 - \alpha)(1 - \alpha^{-1})} - \frac{s^*}{\alpha(1 - \alpha^{-1})} = \frac{1 - \alpha^{s^*}}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{s^*}{1 - \alpha} \end{aligned}$$

and

$$\begin{aligned} J_{42} &= \sum_{m=\nu+1}^{\nu+s^*} \alpha^{-m} \left( \sum_{n=m+1}^{\nu+s^*} \alpha^n \right) \\ &= \sum_{m=\nu+1}^{\nu+s^*} \frac{\alpha^{-m}(\alpha^{m+1} - \alpha^{\nu+s^*+1})}{1 - \alpha} = -\frac{\alpha^{\nu+s^*+1}}{1 - \alpha} \sum_{m=\nu+1}^{\nu+s^*} \alpha^{-m} + \sum_{m=\nu+1}^{\nu+s^*} \frac{\alpha}{1 - \alpha} \\ &= -\frac{\alpha^{\nu+s^*+1}(\alpha^{-\nu-1} - \alpha^{-\nu-s^*-1})}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{\alpha s^*}{1 - \alpha} = -\frac{\alpha^{s^*} - 1}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{\alpha s^*}{1 - \alpha}, \end{aligned}$$

it follows that

$$J_4 = J_{41} + J_{42} = \frac{(1 + \alpha)s^*}{1 - \alpha} + \frac{2(1 - \alpha^{s^*})}{(1 - \alpha)(1 - \alpha^{-1})}.$$

Furthermore,

$$J_1 + J_4 = \frac{4(1 - \alpha^{s^*})}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{2(1 + \alpha)s^*}{1 - \alpha}.$$

Finally, for  $J_2$  and  $J_3$ :

$$\begin{aligned} J_2 &= \left( \sum_{m=\nu-s^*}^{\nu-1} \alpha^{-m} \right) \left( \sum_{n=\nu+1}^{\nu+s^*} \alpha^n \right) \\ &= \left( \frac{\alpha^{-\nu+s^*} - \alpha^{-\nu}}{1 - \alpha^{-1}} \right) \left( \frac{\alpha^{\nu+1} - \alpha^{\nu+s^*+1}}{1 - \alpha} \right) = -\frac{\alpha(\alpha^{s^*} - 1)^2}{(1 - \alpha)(1 - \alpha^{-1})} \end{aligned}$$

and

$$J_3 = \left( \sum_{m=\nu+1}^{\nu+s^*} \alpha^m \right) \left( \sum_{n=\nu-s^*}^{\nu-1} \alpha^{-n} \right) = J_2.$$

We obtain by (2.11) that

$$\begin{aligned} J &= J_1 - J_2 - J_3 + J_4 \\ &= \frac{4(1 - \alpha^{s^*})}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{2\alpha(\alpha^{s^*} - 1)^2}{(1 - \alpha)(1 - \alpha^{-1})} + \frac{2s^*(1 + \alpha)}{1 - \alpha} \\ &= \frac{2s^*(1 - \alpha^2) - 4\alpha(1 - \alpha^{s^*}) - 2\alpha^2(1 - \alpha^{s^*})^2}{(1 - \alpha)^2} =: \sigma_H^2. \end{aligned} \quad (2.12)$$

Next, we compute  $\tilde{J}$ . By (2.4) and (2.10), we can divide  $\tilde{J}$  into four sums:

$$\begin{aligned}\tilde{J} &= \left( \sum_{m,n=\nu-s^*}^{\nu-1} - \sum_{m=\nu-s^*}^{\nu-1} \sum_{n=\nu+1}^{\nu+s^*} - \sum_{m=\nu+1}^{\nu+s^*} \sum_{n=\nu-s^*}^{\nu-1} + \sum_{m,n=\nu+1}^{\nu+s^*} \right) \alpha^{m+n} \\ &= \tilde{J}_1 - \tilde{J}_2 - \tilde{J}_3 + \tilde{J}_4,\end{aligned}$$

where

$$\begin{aligned}\tilde{J}_1 &= \left( \sum_{m=\nu-s^*}^{\nu-1} \alpha^m \right) \left( \sum_{n=\nu-s^*}^{\nu-1} \alpha^n \right) = \left( \frac{\alpha^{\nu-s^*} - \alpha^\nu}{1 - \alpha} \right)^2, \\ \tilde{J}_4 &= \left( \sum_{m=\nu+1}^{\nu+s^*} \alpha^m \right) \left( \sum_{n=\nu+1}^{\nu+s^*} \alpha^n \right) = \left( \frac{\alpha^{\nu+1} - \alpha^{\nu+s^*+1}}{1 - \alpha} \right)^2\end{aligned}$$

and

$$\tilde{J}_2 = \tilde{J}_3 = \left( \sum_{m=\nu-s^*}^{\nu-1} \alpha^m \right) \left( \sum_{n=\nu+1}^{\nu+s^*} \alpha^n \right) = \left( \frac{\alpha^{\nu-s^*} - \alpha^\nu}{1 - \alpha} \right) \left( \frac{\alpha^{\nu+1} - \alpha^{\nu+s^*+1}}{1 - \alpha} \right).$$

Therefore,

$$\begin{aligned}\tilde{J} &= \tilde{J}_1 - \tilde{J}_2 - \tilde{J}_3 + \tilde{J}_4 \\ &= \frac{\alpha^{2\nu}}{(1 - \alpha)^2} ((\alpha^{-s^*} - 1)^2 + \alpha^2(1 - \alpha^{s^*})^2 - 2\alpha(\alpha^{-s^*} - 1)(1 - \alpha^{s^*})) \\ &= \alpha^{2\nu} \frac{(1 - \alpha^{s^*})^2 (\alpha^{-s^*} - \alpha)^2}{(1 - \alpha)^2} =: \tilde{\sigma}_H^2.\end{aligned}\tag{2.13}$$

Finally, by (2.10)–(2.13), we get

$$2s^* \left( \frac{1 - \alpha^2}{\sigma^2} \right) \text{Var}(W_\nu(s)) = J - \tilde{J} = \sigma_H^2 - \tilde{\sigma}_H^2,$$

where  $\sigma_H$  and  $\tilde{\sigma}_H$  are stated in (2.12) and (2.13), respectively. From this, we deduce that the Haar wavelet transform  $W_\nu(s)$  of  $\{x_n\}_0^{N-1}$  is distributed as

$$\frac{\sigma \sqrt{\sigma_H^2 - \tilde{\sigma}_H^2}}{\sqrt{2s^*(1 - \alpha^2)}} X,$$

1 where  $X$  is a normal distribution with mean 0 and variance 1.

2 From this, we get the following theorem.

**Theorem 2.1.** *Let  $\{x_n\}_0^{N-1}$  be an AR(1) time series and  $W_\nu(s)$  be its Haar wavelet transform. Then the Haar wavelet power spectrum of  $\{x_n\}_0^{N-1}$*

$$|W_\nu(s)|^2 \Rightarrow \frac{\sigma^2(\sigma_H^2 - \tilde{\sigma}_H^2)}{2s^*(1 - \alpha^2)} \chi_1^2,$$

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where “ $\Rightarrow$ ” indicates “is distributed as” and  $\chi_1^2$  is the chi-square distribution with one degree of freedom,  $s^* = \frac{s}{2\delta t}$ , and

$$\sigma_H^2 - \tilde{\sigma}_H^2 = \frac{2s^*(1-\alpha^2) - 4\alpha(1-\alpha^{s^*}) - 2\alpha^2(1-\alpha^{s^*})^2}{(1-\alpha)^2} - \alpha^{2\nu} \frac{(1-\alpha^{s^*})^2(\alpha^{-s^*} - \alpha)^2}{(1-\alpha)^2}.$$

**Remark 2.1.** Since  $0 \leq \alpha < 1$ , when  $\nu$  is sufficiently large such that  $\alpha^{2\nu} \approx 0$ , the term

$$\tilde{\sigma}_H \approx 0.$$

The above formula is simplified as follows:

$$|W_\nu(s)|^2 \Rightarrow \frac{\sigma^2 \sigma_H^2}{2s^*(1-\alpha^2)} \chi_1^2,$$

1 where  $\sigma_H^2$  is stated in (2.12).

If  $\{x_n\}_0^{N-1}$  is a white noise with variance  $\sigma^2$ , i.e.  $\alpha = 0$ , then  $\sigma_H^2 - \tilde{\sigma}_H^2 = 2s^*$  and

$$|W_\nu(s)|^2 \Rightarrow \sigma^2 \chi_1^2.$$

### 2 3. Numerical Experiment

3 The AO is a key aspect of climate variability in the Northern Hemisphere. The AO is  
4 defined as the leading empirical orthogonal function (EOF) of Northern Hemisphere

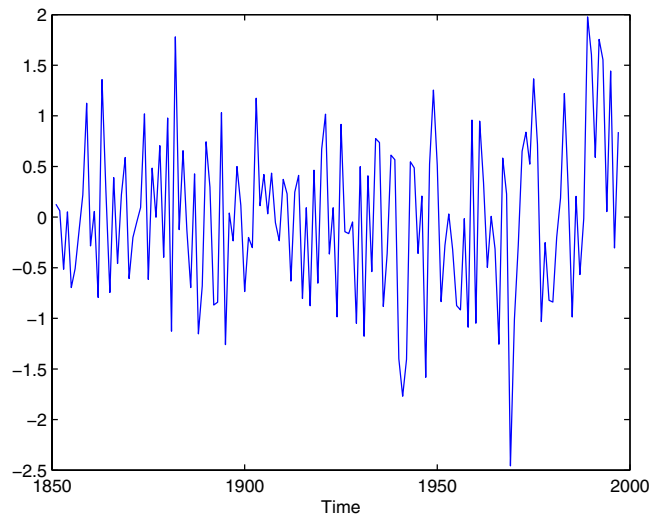


Fig. 1. The standardized time series of winter AO index.



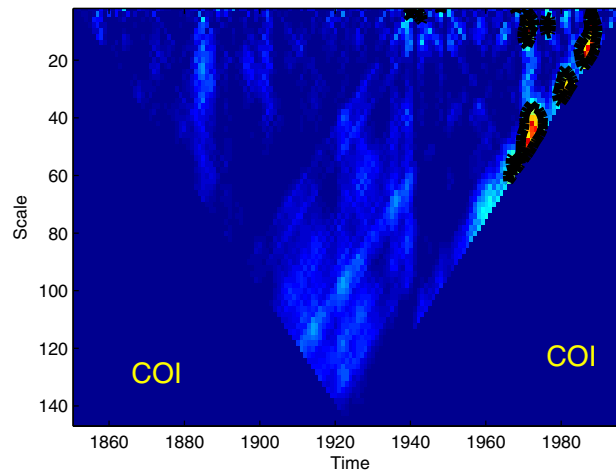
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Fig. 2. Haar wavelet power spectrum of AO index. The thick black contour designates the 5% significance level against red noise. In addition, the COI also be marked (color online).

1 sea level pressure anomalies pole ward of  $20^{\circ}\text{N}$ ,<sup>6</sup> and characterized by an exchange  
 2 of atmospheric mass between the Arctic and middle latitudes. Figure 1 shows the  
 3 winter AO index (December–February 1851–1997). We will use Haar wavelet to  
 4 analyze AO index, i.e. we will compute the Haar wavelet power spectra of the AO  
 5 index. Since the AO index is a finite-length time series, errors in wavelet transform  
 6 will occur at the beginning and end of the wavelet power spectrum. The Cone  
 7 of influence (COI) is the region of the wavelet transform in which edge effects  
 8 become important. Here, we will not research wavelet power spectrum in COI.  
 9 Using Theorem 2.1, we can extract the instinct information of time series from  
 10 background red noise. Figure 2 shows Haar wavelet power spectrum of AO index.  
 11 The thick black contour designates the 5% significance level against red noise. Most  
 12 regions of significance appear abruptly about 1970. Since the Haar wavelet is very  
 13 broadband, this discovers a great change in fundamental properties of the AO,  
 14 commonly called a regime shift or tripping point.

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 20 and Resource Ecology, and Polar Climate and Environment Key Laboratory.

### 21 **Appendix A. Haar Wavelet Scale and Fourier Wavelength**

22 In order to give better physical interpretation, one needs to transfer wavelet scale  
 23 to Fourier wavelength. We will use Meyers' method<sup>5</sup> to obtain the relationship

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1 between the equivalent Fourier period  $T$  and the wavelet scale  $a$ . This relationship  
 2 estimated by computing the wavelet power spectrum of a cosine wave of a known  
 3 frequency and then finding the scale  $a$  at which the wavelet power spectrum reaches  
 4 its maximum. In detail:

Let  $x(t) = \cos \omega t$ . Then

$$\begin{aligned} W(a, b) &= \frac{1}{\sqrt{a}} \int_{-\infty}^{\infty} \cos \omega t H\left(\frac{t-b}{a}\right) dt = \sqrt{a} \int_{-\frac{1}{2}}^{\frac{1}{2}} \cos \omega(at+b) H(t) dt \\ &= -\frac{1}{\omega\sqrt{a}} \left( \sin \omega \left(\frac{a}{2} + b\right) + \sin \omega \left(\frac{a}{2} - b\right) \right) = -\frac{2}{\omega\sqrt{a}} \sin \frac{a\omega}{2} \cos b\omega \end{aligned}$$

and so

$$|W(a, b)|^2 = \frac{4}{\omega^2 a} \sin^2 \frac{a\omega}{2} \cos^2 b\omega.$$

When  $\omega$  is such that  $\sin \frac{\omega a}{2} = 0$  or  $\tan \frac{\omega a}{2} = \omega a$ , we have

$$\frac{d}{da} |W(a, b)|^2 = 0.$$

From this, we deduce that the power spectrum  $|W_\nu(s)|^2$  attains the maximal value at  $a \approx \frac{2.3311}{\omega}$ . Note that  $\omega = \frac{2\pi}{T}$ , where  $T$  is the Fourier wavelength. So we obtain the relationship between the scale parameter and the Fourier wavelength as follows:

$$a \approx \frac{2.3311}{2\pi} T = 0.3710T.$$

5 The relation between scale  $a$  and Fourier period  $T$  for the Morlet wavelet is  
 6  $a = 0.97T$ .<sup>5</sup> However, for Haar wavelet, the corresponding formula is  $a = 0.3710T$ .  
 7 Since for any time series of time step  $\delta t$  and total length  $N\delta t$ , the range of scales  
 8 is from the smallest resolvable scale  $2\delta t$  to the largest scale  $N\delta t$  in wavelet-based  
 9 time series analysis, by using the Haar wavelet analysis one can extract more low  
 10 frequency intrinsic information.

## 11 References

- 12 1. I. Daubechies, The wavelet transform time-frequency localization and signal analysis,  
 13 *IEEE Trans. Inform. Theory* **36** (1990) 961–1004.
- 14 2. A. Grinsted, S. Jevrejeva and J. C. Moore, Application of the cross wavelet trans-  
 15 form and wavelet coherence to geophysical time series, *Nonlinear Process. Geophys.* **11**  
 16 (2004) 561–566.
- 17 3. S. Jevrejeva, J. C. Moore and A. Grinsted, Influence of the arctic oscillation and El  
 18 Nino-southern oscillation (ENSO) on ice conditions in the Baltic sea: The wavelet  
 19 approach, *J. Geophys. Res.* **108** (2003) 4677–4687.
- 20 4. M. E. Mann and J. M. Lees, Robust estimation of background noise and signal detection  
 21 in climatic time series, *Clim. Change* **33** (1996) 409–445.
- 22 5. S. D. Meyers, B. G. Kelly and J. J. O'Brien, An introduction to wavelet analysis in  
 23 oceanography and meteorology with application to the dispersion of Yanai waves, *Mon.*  
 24 *Weather Rev.* **121** (1993) 2858–2866.

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- 1 6. D. W. J. Thompson and J. M. Wallace, The Arctic Oscillation signature in the winter  
2 geopotential height and temperature fields, *Geophys. Res. Lett.* **25** (1998) 1297–1300.
- 3 7. Z. Zhang, Fourier supports of scaling functions determine cardinalities of wavelets, *Int.*  
4 *J. Wavelets Multiresolut. Inf. Process.* **10** (2012) 1–24.
- 5 8. Z. Zhang and J. C. Moore, New significance test methods for Fourier analysis of geo-  
6 physical time series, *Nonlinear Process. Geophys.* **18** (2011) 643–652.